

Mathematical Model of Antibiotic Residues after Applying Poultry Manure to Soil Water



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Abstract

Manure composition emissions can be effected by soil water with the application of poultry manure. The vectors to represent the relationship between soil water characteristics and antibiotic were generated. A Semi-empirical Richards equation is employed to model the law of soil moisture while precise solutions of Richards equation obtained by using the (G'/G) - expansion method and the homogeneous balance method. Then mode I which involves several variables including soil moisture content, soil depth and timescales was assumed according to the exact solutions of Semi-empirical Richards equation. At least one antibiotic was detected in all the soil water and poultry manure samples. And it is worth noting that all of parameters or constants in this function expression were determined by these data.

Keywords: Antibiotic; Poultry manure; Richards Equation; Soil Moisture; The (G'/G) - Expansion Method

Introduction

The large-scale and intensified development of livestock and poultry breeding has resulted in a large number of veterinary antibiotics being used in the aquaculture industry and increasing year by year. Accumulation of antibiotics in the organism. It is very low, after the body absorbs a small part, it undergoes metabolic reactions such as hydroxylation, cleavage, and glucuronidation to produce inactive products, and about 60% to 90% pass through feces as it is. Excretion of urine and excrement of livestock and poultry have undoubtedly become one of the main sources of environmental pollution by antibiotics. The drug design of antibiotics is mainly to kill pathogenic pathogens. Once they enter the environment, they will inevitably pose a potential threat to other organisms in the environment. Firstly, the resistant pathogenic bacteria or variant pathogens in livestock and poultry are produced and continuously released into the environment. Secondly, livestock and poultry continue to excrete these antibiotics or their metabolites into the environment, so that drug resistant pathogens and mutant pathogens in the environment are continuously generated.

Both of these in turn stimulated the producers to increase the dosage and renew the drug varieties, which resulted in a vicious circle of "drug-contaminated environment → emergence of drug-resistant or mutated pathogens → increased dosages → environmental pollution". Fertilization by poultry manure has

shown an important variance in soil water chemical characteristics. The state for regional soil moisture reserve is the strategic storage of water resources in the district. The distribution of soil moisture directly affects the supply of groundwater resources, determines the amount of water, which is absorbed from the soil and evaporated by the earth's surface plants, plays a decisive factor for plants' productivity, and also is regarded as strategic factors influencing the ecological environment security, the economic development and the people's lives in arid and semi-arid areas [1].

Presence of antibiotic and soil water physicochemical properties played key roles in degradation of numerous molecules and other processing. Fertilization is the commonest managing agricultural soils, and for a long time, intensive farming appealed to fertilizer to increase yields. Then livestock and poultry excrement will constitute one of the main reasons for the non-point source pollution of antibiotics in China. Therefore, based on the research results in recent years, combining the use of antibiotics in the livestock and poultry industry and the status of residues, this paper analyzes the fate and environmental risks of livestock and poultry excrement in soil, and puts forward the corresponding mathematical model. Darcy law is a fundamental theoretical method to describe the motion law of soil moisture, therefore a variety of Richards equations are deduced. For the nonlinear

partial differential equations, the previous research method is to discuss their definite solutions and we commonly can acquire their numerical solution through the numerical method.

Whether the analytical solution of Richards's equations, which describes the change of soil moisture content with the change of time and space position in Darcy's law, has been the expectation, If we substitute some empirical representations of hydraulic conductivity and water diffusivity into Richards equations, the exact solution of Richards's equation on soil moisture content, soil depth and time is of great significance. Furthermore, the greater parts of antibiotic were found in the soil water by poultry manure. In the past few years, many powerful methods to construct exact solutions of nonlinear evolution equations have been established and developed such as the homogeneous balance method [2-4], the (G'/G) -expansion method [5,6], the exp-function method [7,8] and so on.

One of the most effective and direct methods for constructing exact solutions of nonlinear differential equations is the (G'/G) - expansion method. The (G'/G) - expansion method, first introduced by Wang et al. [5], has been widely used to search for various exact solutions of NLEEs [9-11]. The (G'/G) - expansion method is based on the [12,13] explicit linearization of nonlinear differential equations for traveling waves with a certain substitution which leads to a second-order differential equation with constant coefficients [14-16]. Finding an exact solution for Richards equation, by using the (G'/G) - expansion method, is the main goal of the present study.

Mathematical Models and Explanations

First of all, we introduce a form of Richards's equations as follows:

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial x} \left[D(\theta) \frac{\partial \theta}{\partial x} \right] + \frac{\partial}{\partial y} \left[D(\theta) \frac{\partial \theta}{\partial y} \right] + \frac{\partial}{\partial z} \left[D(\theta) \frac{\partial \theta}{\partial z} \right] + \frac{\partial K(\theta)}{\partial z}, \tag{2.1}$$

where $D(\theta)$ denotes water diffusivity; $K(\theta)$ denotes hydraulic conductivity; t denotes time; θ denotes soil moisture content; x, y, z denote coordinate axes. If the soil moisture content is lower than the saturated (unsaturated) moisture content with little change, we take as $D(\theta) = a$, where a is a constant.

Many researchers have committed themselves to estimating soil hydraulic conductivity, as a result, various empirical representations of hydraulic conductivity are proposed. We assume that [12,13] unsaturated hydraulic conductivity is calculated by using the Libardi method, that is

$$K(\theta) = K_0 \exp\{\beta(\theta - \theta_0)\} \tag{2.2}$$

Where β is a constant; K_0 and θ_0 are the values of K and θ during steady-state infiltration, respectively.

Next, we have intend to simplify equation (2.1), in other words, here we only consider the case that soil moisture flows in the vertical direction, and therefore we have

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial z} \left[D(\theta) \frac{\partial \theta}{\partial z} \right] + \frac{\partial K(\theta)}{\partial z}. \tag{2.3}$$

By substituting $D(\theta) = a$ and equation (2.2) into equation (2.3),

hence the following semi-empirical Richards equation is obtained:

$$\frac{\partial \theta}{\partial t} = a \frac{\partial^2 \theta}{\partial z^2} + K_0 \beta \exp\{\beta(\theta - \theta_0)\} \frac{\partial \theta}{\partial z}. \tag{2.4}$$

In this section, by make use of the (G'/G) -expansion method, we obtain an exact solution for the equation (2.4), however, we omit the description of the (G'/G) -expansion method. If you are interested in this method, you can refer to the reference [4].

Let $\frac{\partial \theta}{\partial t} = \theta_t$, $\frac{\partial \theta}{\partial z} = \theta_z$ and $\alpha = K_0 \exp\{-\beta\theta_0\}$, then the equation (2.4) can be equivalently changed into

$$a\theta_{zz} + \alpha\beta e^{\beta\theta} \theta_z - \theta_t = 0. \tag{2.5}$$

Using the travelling wave variable $\theta(z, t) = \theta(\xi)$ and $\xi = z - \omega t$ carries out the equation (2.5) into an ordinary differential equation for $\theta = \theta(\xi)$

$$a\theta'' + \alpha\beta e^{\beta\theta} \theta' + \omega\theta' = 0. \tag{2.6}$$

In order to apply the (G'/G) -expansion method, we use the Painlevé transformation $v = e^{\beta\theta}$, or equivalently $\theta = \frac{1}{\beta} \ln v$, hence the equation (2.6) can be written as

$$a(vv'' - v'v') + \alpha\beta v'v^2 + \omega v'v = 0. \tag{2.7}$$

Suppose that the solution of ordinary differential equation (2.7) can be expressed by a polynomial in (G'/G) as follows:

$$v(\xi) = \sum_{i=1}^n a_i \left(\frac{G'}{G}\right)^i + a_0, \quad a_n \neq 0, \tag{2.8}$$

where $G = G(\xi)$ satisfies the second order LODE in the form

$$G'' + \lambda G' + \mu G = 0 \tag{2.9}$$

where $G' = \frac{dG}{d\xi}$, $G'' = \frac{d^2G}{d\xi^2}$, $a_n, \dots, a_0, \lambda, \mu$ are real constants to be determined later.

According to the (G'/G) - expansion method, considering the homogeneous balance between vv'' and $v'v^2$ in the equation (2.7), we get $3n + 1 = 2n + 2 \Rightarrow n = 1$, hence we can write (2.8) as

$$v(\xi) = a_1 \left(\frac{G'}{G}\right) + a_0, \quad a_1 \neq 0. \tag{2.10}$$

Substituting (2.10) along with (2.9) into (2.7) and collecting all terms with the same order of $\left(\frac{G'}{G}\right)$ together, the left-hand side of (2.7) are converted into polynomial in $\left(\frac{G'}{G}\right)$. Setting each coefficient of each polynomial to zero, we derive a set of algebraic equations for $\lambda, \mu, \omega, a_0, a_1$ as follows:

$$\begin{aligned} \left(\frac{G'}{G}\right)^4 &: aa_1^2 - \alpha\beta a_1^3, \\ \left(\frac{G'}{G}\right)^3 &: -\omega a_1^2, \\ \left(\frac{G'}{G}\right)^2 &: 3a\lambda a_0 a_1 - \alpha\beta(a_1 a_0^2 + \mu a_1^3 + 2\lambda a_0 a_1^2) - \omega(a_1 a_0 + \lambda a_1^2), \\ \left(\frac{G'}{G}\right)^1 &: a(2\mu + \lambda^2)a_0 a_1 - a\lambda \mu a_1^2 - \alpha\beta(\lambda a_1 a_0^2 + 2\mu a_0 a_1^2) - \omega(\lambda a_1 a_0 + \mu a_1^2), \\ \left(\frac{G'}{G}\right)^0 &: a(\lambda \mu a_0 a_1 - \mu^2 a_1^2) - \alpha\beta(\mu a_1 a_0^2) - \omega \mu a_1 a_0. \end{aligned}$$

Solving the algebraic equations above yields

$$a_1 = \frac{a}{\alpha\beta}, \quad \omega = 0, \quad \lambda = \frac{a_0 \alpha \beta}{a}, \quad \mu = 0, \tag{2.11}$$

where a_0 is arbitrary constant.

By using (2.11), (2.10) can be written as

$$v(\xi) = \frac{a}{\alpha\beta} \left(\frac{G'}{G} \right) + a_0 \tag{2.12}$$

where $\xi = z$.

Substituting the general solutions of equation (2.9) into (2.12), we have an exact solution of the equation (2.7) as follows:

$$v(\xi) = \frac{a}{\alpha\beta} \frac{\sqrt{\lambda^2 - 4\mu}}{2} \left(\frac{C_1 \sinh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \xi + C_2 \cosh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \xi}{C_1 \cosh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \xi + C_2 \sinh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \xi} \right) + \frac{a_0}{2} \tag{2.13}$$

where $\xi = z$, $\lambda^2 - 4\mu = \left(\frac{a_0}{\alpha} \alpha\beta \right)^2$, $\alpha = K_0 \exp\{-\beta\theta_0\}$, K_0 , θ_0 , C_1 , C_2 , a , β and a_0 are arbitrary constants.

Therefore by $\theta = \frac{1}{\beta} \ln v$, we have an exact solution of the equation (2.4) as follows:

$$\theta = \frac{1}{\beta} \ln \left\{ \frac{a}{\alpha\beta} \frac{\sqrt{\lambda^2 - 4\mu}}{2} \left(\frac{C_1 \sinh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \xi + C_2 \cosh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \xi}{C_1 \cosh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \xi + C_2 \sinh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \xi} \right) + \frac{a_0}{2} \right\} \tag{2.14}$$

According to what has been discussed above, we assume that soil moisture content, soil depth and time satisfy

$$\theta = \frac{1}{\beta} \ln \left\{ \frac{a}{\alpha\beta} \frac{\sqrt{\lambda^2 - 4\mu}}{2} \left(\frac{C_1 \sinh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \zeta + C_2 \cosh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \zeta}{C_1 \cosh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \zeta + C_2 \sinh \frac{\sqrt{\lambda^2 - 4\mu}}{2} \zeta} \right) + \frac{a_0}{2} \right\} \tag{2.15}$$

where $\zeta = \delta z - \epsilon t$, δ , ϵ are arbitrary constants.

Materials and Methods

Data used here were collected during an internal drainage experiment, carried out on a sandy-loam Red Yellow Latosol (Typic Hapludox) of the county of Piracicaba, SP, Brazil, with a fairly homogeneous profile down to the depth of 2m [17]. They consist of soil water content values calculated from tensiometer readings, through the use of laboratory established soil water retention curves and of soil water potential heads, measured with the same mercury manometer tensiometers. The soil water retention curves were established by drying wet samples, to minimise hysteresis effects. Data are averages of 3 replicates collected during the internal drainage of three 3 m by 3m plots, 10m apart, previously submitted to infiltration by ponding water on the soil surface until quasi-steady flow. To prevent evaporation losses, the soil surface

Table 1: The soil profile characterization.

Granulometric analysis									
Soil fraction	Depth (m)								
	0.15	0.3	0.45	0.6	0.75	0.9	1.05	1.2	1.35
Clay (%)A	18.17	27.88	26.82	27.1	28.11	27.87	26.41	28.31	26.78
Silt (%)	4.72	4.22	5.29	3.93	3.08	3.16	5.51	3.61	5.2
Sand (%)	77.11	67.9	67.89	68.97	68.81	68.97	68.87	68.08	68.02
Bulk density									
(kg/m ³)	1510	1440	1400	1350	1350	1360	1410	1400	1370
Soil-water retention curve									
Matric potential									
Head (mH ₂ O)									

was covered with a plastic sheet during the drainage process, so that all conditions stated in Libardi et al. [12] were met [18]. Soil-water retention curve at each soil depth was determined using undisturbed samples (45mm diameter by 45 mm height) on porous plate funnels (matric potential head h from 0.26m to 1.87m) and porous plate pressure cells (matric potential head h from 3.0 m to 8.0m).

Soil bulk densities were also determined using volumetric cylinders of 45mm diameter and 45mm height. Table 2 also presents the obtained values of parameters a, m, n, r, and the equation of Van Genuchten [19] as well as their r² values; soil water content values during the redistribution time were obtained by means of these equations. The saturated soil water contents θ were also taken from the Van Genuchten equation for matric potential h=0. The determination of θ in the field at steady infiltration is practically impossible, except with neutron probes, which have to be well calibrated and present difficulties for measurements close to the soil surface. It is important to note that it is difficult to fully saturate a soil profile in the field and that; therefore, field values of θ are, in general, slightly lower than those measured in the laboratory. Tables 2 and 3 present the average raw data of soil water content and matric potential head, which are presented here for any other development of soil water movement theories. Soil water potential head H (cm of water) is considered as the sum of the gravitational head z and the matric potential head h [17].

Results and Discussion

In this model, the data were collected during an internal drainage experiment, which was carried out on a sandy-loam Red Yellow Latosol (Typic Hapludox) of the county of Piracicaba [17]. We use the data to determine the parameters in the identity (2.15), and therefore the predicted values of soil moisture content are calculated from (2.15). From data in (Tables 1&2), it revealed that volumetric soil water contents (m³/m³), for different redistribution times (t) and different depths (z). And the parameters of equation (2.15) for each time and soil depth were determined by using data fitting method. All of these can be seen in Table 1. These parameters are shown in Table 3. From these parameters, we obtain a function express in which soil water content is independent variable whereas time and soil depth are two dependent variables.

-0.00B	0.43	0.457	0.472	0.491	0.491	0.487	0.468	0.472	0.483
-0.26	0.29	0.35	0.344	0.377	0.335	0.325	0.329	0.352	0.348
-0.46	0.25	0.32	0.305	0.333	0.286	0.289	0.291	0.312	0.305
-0.97	0.22	0.28	0.257	0.271	0.219	0.227	0.231	239	0.236
-1.41	0.2	0.25	0.239	0.23	0.193	0.206	0.207	0.211	0.209
-1.87	0.18	0.23	0.215	0.208	0.177	0.192	0.194	0.196	0.194
-3.00	0.167	0.201	0.19	0.173	0.172	0.173	0.177	0.171	0.161
-5.00	0.148	0.188	0.175	0.159	0.157	0.157	0.162	0.155	0.149
-8.00	0.133	0.178	0.168	0.15	0.148	0.148	0.152	0.145	0.138

Table 2: Volumetric soil water contents θ (m^3/m^3), for different redistribution times (t) and different depths (z).

Time (h)	0.15	0.3	0.45	Soil depth (m)		0.9	1.05	1.2	1.35
				0.60	0.75				
0.25	0.383	0.418	0.41	0.459	0.453	0.41	0.405	0.425	0.427
0.5	0.356	0.407	0.393	0.444	0.43	0.386	0.384	0.41	0.413
0.75	0.337	0.399	0.381	0.435	0.411	0.371	0.367	0.397	0.401
1	0.323	0.392	0.372	0.427	0.397	0.357	0.354	0.385	0.391
1.25	0.314	0.385	0.363	0.42	0.385	0.347	0.343	0.376	0.383
1.5	0.307	0.38	0.357	0.413	0.372	0.338	0.333	0.368	0.375
1.75	0.301	0.376	0.351	0.407	0.363	0.33	0.326	0.36	0.369
2	0.297	0.372	0.345	0.402	0.354	0.323	0.32	0.352	0.363
2.5	0.289	0.366	0.337	0.393	0.34	0.312	0.31	0.341	0.353
3	0.283	0.361	0.331	0.385	0.328	0.303	0.302	0.333	0.344
4	0.276	0.354	0.322	0.371	0.311	0.292	0.291	0.32	0.331
5	0.27	0.349	0.314	0.363	0.3	0.285	0.283	0.312	0.319

Table 3: Parameters of equation (2.15).

a	κ_0	β	a_0	C_1	C_2	ϵ	θ_0	δ
1.391	0.627	1.114	1.027	0.513	0.626	-1.066	0.493	-0.308

And thus the predicted values of soil moisture content are also presented in (Table 4) according to equation (2.15). Compared with the actual values, the mean squared error (MSE) is 0.00077423, which suggests that equation (2.15) describes the changes of soil water content with the time and soil depth very well. According to Table 2, the following Figure 1 can be presented. Figure 2 show that

Table 4: The predicted values.

Time (h)	0.15	0.3	0.45	Soil depth (m)		0.9	1.05	1.2	1.35
				0.60	0.75				
0.25	0.408	0.41	0.413	0.416	0.419	0.421	0.424	0.427	0.431
0.5	0.394	0.397	0.399	0.401	0.404	0.406	0.409	0.411	0.414
0.75	0.382	0.384	0.386	0.388	0.39	0.393	0.395	0.397	0.399
1	0.372	0.374	0.375	0.377	0.379	0.381	0.383	0.385	0.387
1.25	0.363	0.365	0.366	0.368	0.369	0.371	0.372	0.374	0.376
1.5	0.355	0.357	0.358	0.359	0.361	0.362	0.363	0.365	0.366
1.75	0.349	0.35	0.351	0.352	0.353	0.354	0.356	0.357	0.358
2	0.343	0.344	0.345	0.346	0.347	0.348	0.349	0.35	0.351
2.5	0.333	0.334	0.335	0.336	0.336	0.337	0.338	0.339	0.34
3	0.326	0.327	0.327	0.328	0.328	0.329	0.33	0.33	0.331

equation (2.15) has the same tendency with the Figure 1. Although the contents are not equal at the range of 0-14m, it can be recorded as the same contents duo to little change in this range. Since the hypothesis that the water diffusivity is constant, Figure 2 obviously conforms to the actual situation.

4	0.317	0.317	0.317	0.318	0.318	0.318	0.319	0.319	0.319
5	0.311	0.311	0.312	0.312	0.312	0.312	0.312	0.313	0.313

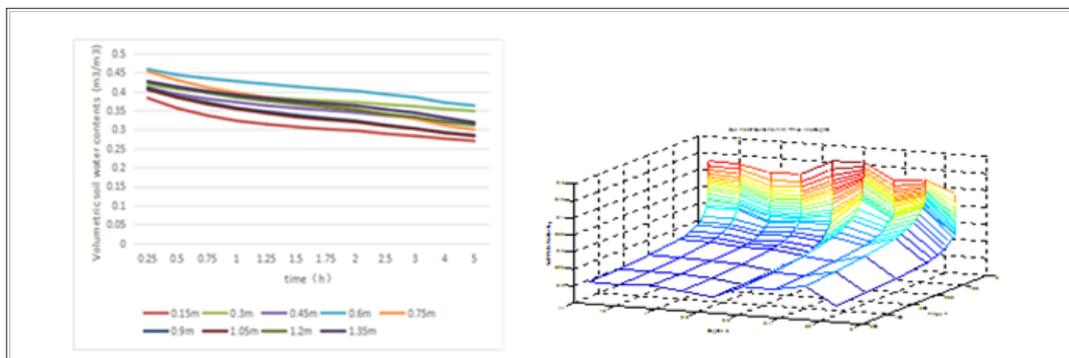


Figure 1: Original figure.

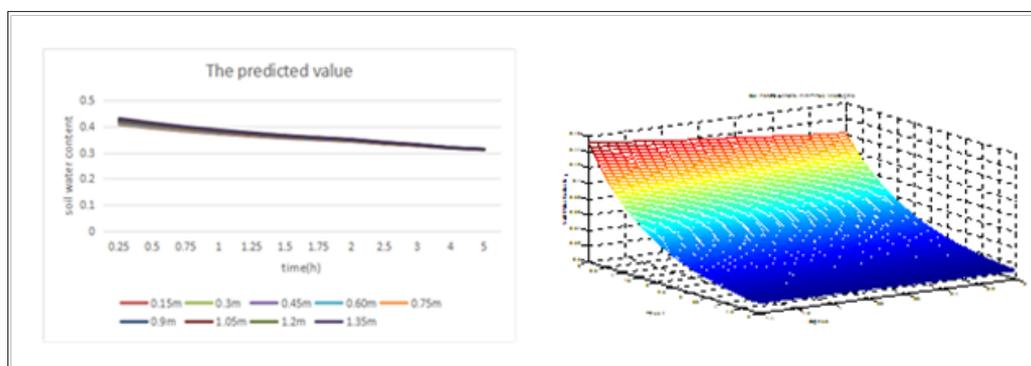


Figure 2: The smoothed figure.

Conclusion

We propose a new model equation (2.15), which describes the changes of soil water content with the time and soil depth, by finding out an exact solution of equation (2.4) through currently prevalent (G' / G) -expansion method. Even though equation (2.15) is established with the assumptions that the hydraulic conductivity satisfies exponential function and water diffusivity is a constant, to some extent, the very low mean squared error (MSE) indicates that equation (2.15) is a great reflection in change of soil water content with the time and soil depth. However, the simple assumption of water diffusivity might incur larger error in term of actual values of soil water content, thus a precise empirical representation of water diffusivity is expected to be proposed. With the rapid development of science and technology, the solutions to nonlinear differential equations would be enriched, and more excellent results on Darcy’s law might be acquired. All these results also suggest that soil water play a considerable role in the fate of antibiotic in the environment. Further research is needed to give the detection of antibiotic from animal manure in soil water by the regression equations constructed in the present study.

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